



Aotearoa New Zealand Initial Ceilometer Analysis

Prepared for University of Canterbury

June 2024

Prepared by:
Daniel Morrish




For any information regarding this report please contact:

Daniel Morrish
Air Quality Technician
Water and Air Quality Impacts
+64-9-375-2057
Daniel.Morrish@niwa.co.nz

National Institute of Water & Atmospheric Research Ltd
Private Bag 99940
Viaduct Harbour
Auckland 1010

Phone +64 9 375 2050

NIWA CLIENT REPORT No: 2024209AK
Report date: June 2024
NIWA Project: UOC23101

Quality Assurance Statement		
	Reviewed by:	Elizabeth Somervell Air Quality Scientist Urban Air Quality and Health
	Formatting checked by:	Elizabeth Somervell Air Quality Scientist Urban Air Quality and Health
	Approved for release by:	Jonathan Moores Regional Manager Auckland

© All rights reserved. This publication may not be reproduced or copied in any form without the permission of the copyright owner(s). Such permission is only to be given in accordance with the terms of the client's contract with NIWA. This copyright extends to all forms of copying and any storage of material in any kind of information retrieval system.

Whilst NIWA has used all reasonable endeavours to ensure that the information contained in this document is accurate, NIWA does not give any express or implied warranty as to the completeness of the information contained herein, or that it will be suitable for any purpose(s) other than those specifically contemplated during the project or agreed by NIWA and the client.

Contents

- Executive summary4**

- 1 Introduction5**

- 2 Methods.....6**
 - 2.1 Data6
 - 2.2 Post processing7
 - 2.3 Threshold Selection.....8
 - 2.4 Rainfall.....9

- 3 Climatologies.....9**

- 4 Further Analysis10**

- 5 Open Data11**

- 6 Acknowledgements.....11**

- 7 References.....12**

- Appendix A Cloud occurrence as a function of altitude by hour13**

Tables

- Table 1: Details of Aotearoa New Zealand Metservice Ceilometers with available data. 7

Figures

- Figure 1 Locations of Aotearoa New Zealand Metservice Ceilometers with available data. 6
- Figure 2 Detail of ALCF output plot showing the cloud mask in red for default and selected cloud mask thresholds. 8
- Figure 3 Cloud occurrence plots from ALCF output of default and selected cloud mask thresholds at locations with available data. 9
- Figure 4 ALCF backscatter plot at Auckland Airport during heavy rainfall January 27th, 2023. 9

Executive summary

Cloud cover significantly impacts the Earth's energy balance by reflecting incoming solar energy and trapping infrared radiation. Accurate representation of clouds in climate models is crucial for predicting climate change. However, biases in cloud simulations remain, particularly in high latitudes, as evidenced by the World Climate Research Programme's CMIP3 models. These biases often stem from inadequate cloud cover representation and difficulties in simulating super-cooled liquid water clouds, leading to errors in shortwave radiation estimates. This study aims to address these biases by analysing cloud data from ceilometers across Aotearoa New Zealand, focusing on the proportion of super-cooled liquid water clouds.

Data was collected from a network of 16 Vaisala CL31 ceilometers operated by MetService at various airports across Aotearoa New Zealand, covering both North and South Islands and Chatham Island. The ceilometers, primarily used for detecting cloud base height and vertical visibility, provided data from July 2021 to February 2023. The uncalibrated backscatter data was converted to NetCDF format using `cl2nc` software and processed with the Automatic Lidar and Ceilometer Framework (ALCF) for consistent noise removal, calibration, and cloud detection.

The default backscatter threshold for cloud detection, developed using LIDAR data from Antarctica, was found to be too low for mid-latitudes, leading to misclassification of surface aerosols and instrument artifacts as clouds. A higher threshold was selected to improve cloud detection accuracy. This adjustment significantly reduced the misclassification of low-altitude clouds while maintaining the detection of higher altitude clouds.

Rainfall, distinguishable in backscatter plots as diagonal bands, can interfere with cloud detection. The backscatter of rain can sometimes be as strong as clouds, complicating the accuracy of the automatic cloud identification mask generated by ALCF. Future analyses will incorporate rain gauge data to filter out periods of rainfall more effectively.

Hourly climatologies of cloud occurrence showed minimal variation between sites, with a slight increase in higher altitude clouds between 1100 and 1700 at most Te Ika-a-Māui / North Island sites. Most sites had a maximum cloud base height of around 2000 metres although some were higher at around 2500 metres, and some lower around 1500 metres. This could be due to weather patterns or surrounding terrain; further investigation is required. The available data was insufficient for developing monthly climatologies.

Future steps will include developing a rain mask, dividing the data based on the prevailing south westerly wind patterns and further analysing cloud heights concerning wind direction. Advanced models capable of classifying super-cooled liquid water clouds, such as G22-Davis and G22-Christchurch, will be used to refine climatologies and improve the accuracy of cloud representation in climate models.

1 Introduction

The Earth energy balance is influenced by cloud cover, affecting surface temperature by reflecting incoming solar energy and trapping infrared radiation in the atmosphere. Observation and model simulation comparisons reveal biases in cloud representation. Large biases were noticeably found over high latitudes in The World Climate Research Programme (WCRP) Coupled Model Intercomparison Project CMIP3 models (Trenberth & Fasullo, 2010). Although further research has improved cloud simulations and their properties, biases remain and can include compensating errors that hide the bias (Schuddeboom & McDonald, 2021; Kuma et al., 2023). Prior research has found that biases are contributed to by insufficient cloud cover and issues with the amount of super-cooled water clouds simulated. The latter problem arises because different types of clouds influence radiation differently. In particular, models often struggle to accurately simulate super-cooled liquid water clouds, leading to significant biases in shortwave radiation (Bodas-Salcedo et al., 2016; Kay et al., 2016; Kuma, 2020). These clouds, which form between the 0°C isotherm and the -38°C isotherm (the homogeneous freezing level), are prevalent over the Southern Hemisphere, especially over the Southern Ocean (Bodas-Salcedo et al., 2016; Kuma, 2020) and Antarctica (Listowski et al., 2019) For example, (Listowski et al., 2019) found that the fraction of super-cooled liquid-water-containing clouds (SLC) ranged from 0-35% over the Antarctic continent. These issues are significant because changes in the global Effective Climate Sensitivity (ECS) between CMIP5 and CMIP6 have been largely attributed to changes in the representation of extra-tropical Southern Hemisphere clouds (Zelinka et al., 2020).

This analysis provides a first look at cloud data from a network of ceilometers across Aotearoa New Zealand and details some of the challenges in processing the data available. The initial analysis provides a better understanding of the data to be used in preparation for following previous methods to determine levels of supercooled liquid water clouds across Aotearoa New Zealand.

2 Methods

2.1 Data

A network of Vaisala CL31 ceilometers operated by MetService at airports across Aotearoa New Zealand had data available for analysis. Nine sites were in Te Ika-a-Māui / North Island, six in Te Waipounamu / South Island and one site on Chatham Island. The 16 sites are shown in figure 1 and details in table 1.



Figure 1 Locations of Aotearoa New Zealand MetService Ceilometers with available data.

Table 1: Details of Aotearoa New Zealand Metservice Ceilometers with available data.

SITE	Lon	Lat	Altitude	Viewing Angle
Kaitaia Aero (NZKTA)	173.2863	-35.0678	76 m	zenith
Whangarei Aero (NZWRA)	174.3655	-35.7669	37 m	zenith
Auckland Aero (NZAAA)	174.8072	-37.009	6 m	zenith
Hamilton Aero (NZHNA)	175.3318	-37.8609	52 m	zenith
Gisborne Aero (NZGSA)	177.9851	-38.6586	4 m	zenith
New Plymouth Aero (NZNPA)	174.1838	-39.0079	27 m	zenith
Napier Aero (NZNRA)	176.8646	-39.4703	2 m	zenith
Ohakea Aero (NZOHA)	175.372	-40.203	46 m	zenith
Nelson Aero (NZNSA)	173.2161	-41.3008	5 m	zenith
Wellington Aero (NZWNA)	174.8057	-41.3312	3 m	12°
Hokitika Aero (NZHKA)	170.9844	-42.713	39 m	zenith
Christchurch Aero (NZCHA)	172.5276	-43.4893	35 m	zenith
Chatham Island Aero (NZCIA)	-176.475	-43.8168	11 m	zenith
Queenstown Aero (NZQNA)	168.7402	-45.0176	356 m	zenith
Dunedin Aero (NZDNA)	170.1968	-45.9268	1 m	zenith
Invercargill Aero (NZNVA)	168.3178	-46.411	2 m	zenith

CL31 ceilometers have a maximum vertical range of 7.6 km, a measurement resolution of 10 m and a measurement interval of 2 s. and are primarily for airport operations detecting cloud base height and vertical visibility. Data was available between and July 2021 to February 2023. The uncalibrated data output files of attenuated volume backscattering coefficients were converted to NetCDF using `cl2nc` (Kuma, 2017/2023).

2.2 Post processing

Post processing was done using the Linux software package: Automatic Lidar and Ceilometer Framework (ALCF). ALCF provides consistent processing of the data by applying the same level of noise removal, calibration, time resampling, height resampling, cloud detection and cloud base detection to each of the instruments (Kuma et al., 2021).

2.3 Threshold Selection

An automatic cloud detection mask using the ceilometer backscatter values is applied within the ALCF processing. The default backscatter threshold for cloud detection of $2 \times 10^{-6} \text{m}^{-1} \text{sr}^{-1}$ was developed using LIDAR data from Antarctica. This value has been too low in previous uses of the tool in Aotearoa New Zealand (Whitehead et al., 2023) and potentially misclassifies surface aerosols and instrument artefacts in the mid-latitudes. Figure 2 shows the difference in how the threshold is applied in the backscatter plots. The cloud mask, indicated by the red line, with the default threshold encompasses much of the weaker 'haze' beneath the clouds, while the cloud mask with the stronger threshold ($6 \times 10^{-6} \text{m}^{-1} \text{sr}^{-1}$) envelopes the cloud while excluding much of the haze. This can be seen in Figure 3 with over 50% of cloud occurrence below 500m at some sites across Aotearoa New Zealand using the default threshold. A stronger threshold of $6 \times 10^{-6} \text{m}^{-1} \text{sr}^{-1}$ was selected for further analysis to reduce the low-altitude mis-classified clouds while not greatly affecting higher altitude cloud detection. Increasing the cloud mask threshold above this provided diminishing returns in classifying the clouds.

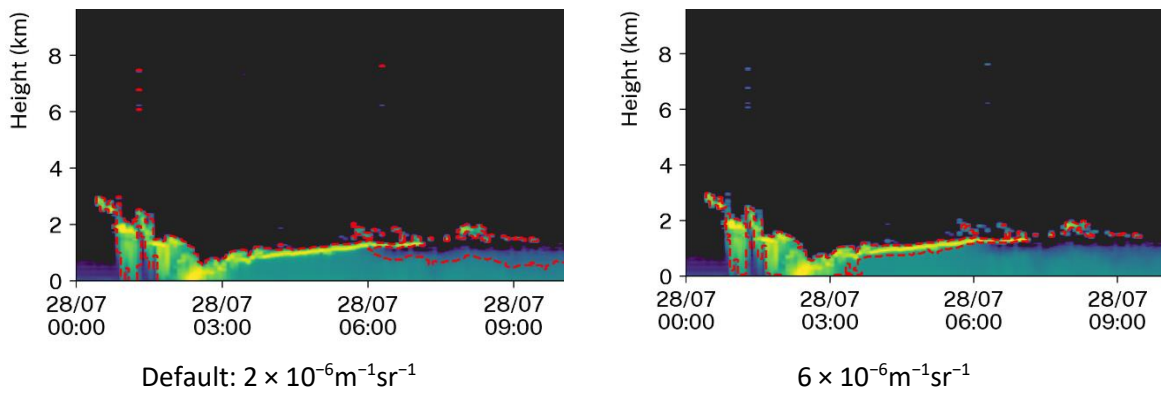


Figure 2 Detail of ALCF output plot showing the cloud mask in red for default and selected cloud mask thresholds.

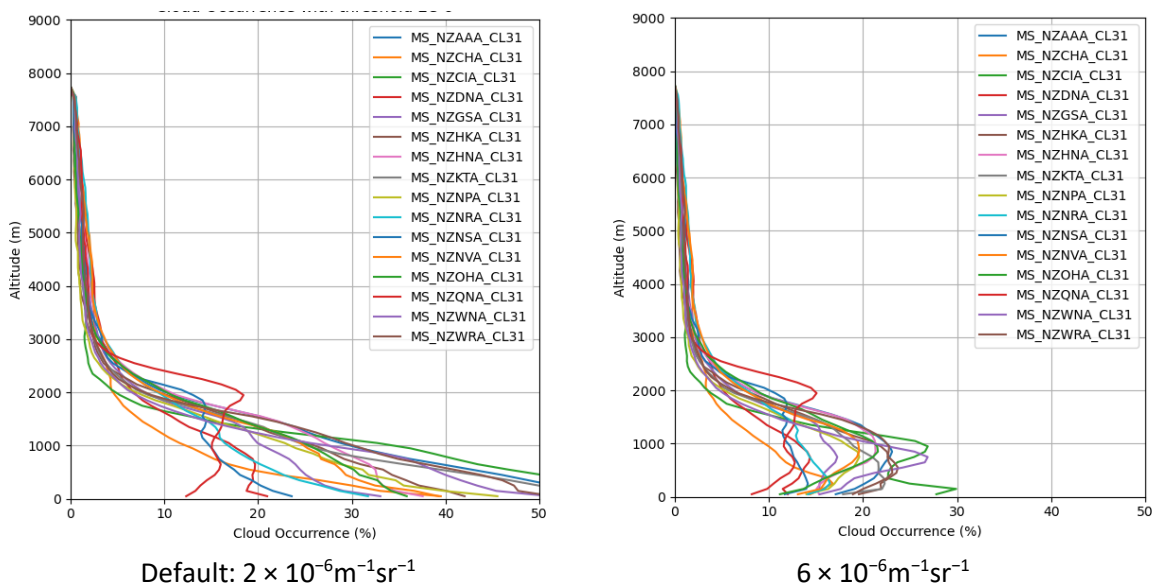


Figure 3 Cloud occurrence plots from ALCF output of default and selected cloud mask thresholds at locations with available data.

2.4 Rainfall

Rainfall can often be easily identified from backscatter plots, as moving diagonal bands coming from a band of stronger signal with a band of cloud above. The backscatter of rain can sometimes be as strong as clouds and can be picked up in the cloud mask (Figure 4). It is not possible to change the cloud mask threshold to remove the mis-classified rain as the cloud mask is created from the strength of the backscatter, and so other methods will need to be used to filter out the times with rainfall such as using rain gauge data from the ceilometer locations to create a rain mask.

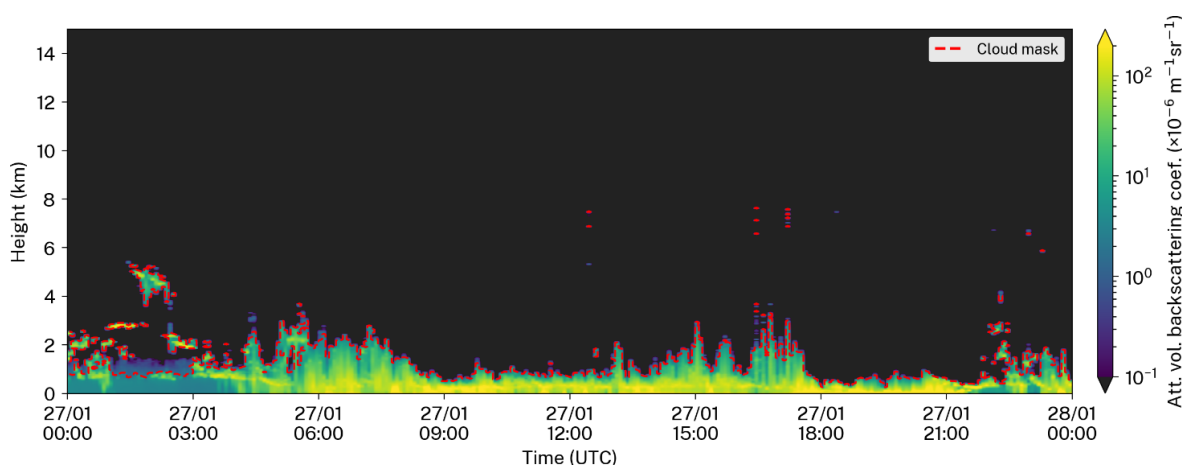


Figure 4 ALCF backscatter plot at Auckland Airport during heavy rainfall January 27th, 2023.

3 Climatologies

Hourly climatologies of cloud occurrence showed little variation between each site, with similar patterns throughout the day. Variation across different hours of higher altitude cloud occurred between 1100 and 1700 hours at all Te Ika-a-Māui / North Island sites except Wellington. Nelson, Napier, Ohakea, Queenstown had a higher maximum cloud base height around 2500 metres. The higher maximum could be due to the higher surrounding terrain, although this would not explain Ohakea, and other sites also have higher surrounding terrain so may have other influences. Wellington, Christchurch, Chatham Islands had a lower maximum cloud base height of around 1500 metres. This could be due to the wind patterns at these locations, but this could also be an influence at other sites and so further investigation is required. All other sites had a maximum cloud base height of around 2000 metres. Queenstown was the only site that had visually distinct cloud layers compared to all other sites, potentially due to the site's location in the Southern Alps. Not enough data was available for monthly cloud occurrence climatologies. Hourly climatologies of cloud height at each site are presented in Appendix A.

4 Further Analysis

Further steps should include applying a rain mask to remove the , analysing a south-westerly transect of the sites following the country's prevailing winds. Also, a further split of wind direction at each site to explore variation of cloud heights should be conducted. Climatologies should then be developed when the data has been passed through a model that can classify sections of the clouds that contain super cooled liquid water such as G22-Davis (Guyot et al., 2022) G22-Christchurch (Whitehead et al., 2023).

5 Open Data

Level 0 (raw) ceilometer data is owned by MetService and is available with permission from MetService. Level 1 (ALCF output) is available as a processed data product and is available at Zenodo. <https://doi.org/10.5281/zenodo.12556402>

6 Acknowledgements

This study was funded by the Deep South National Science Challenge. NIWA thanks MetService for providing access to Ceilometer Data.

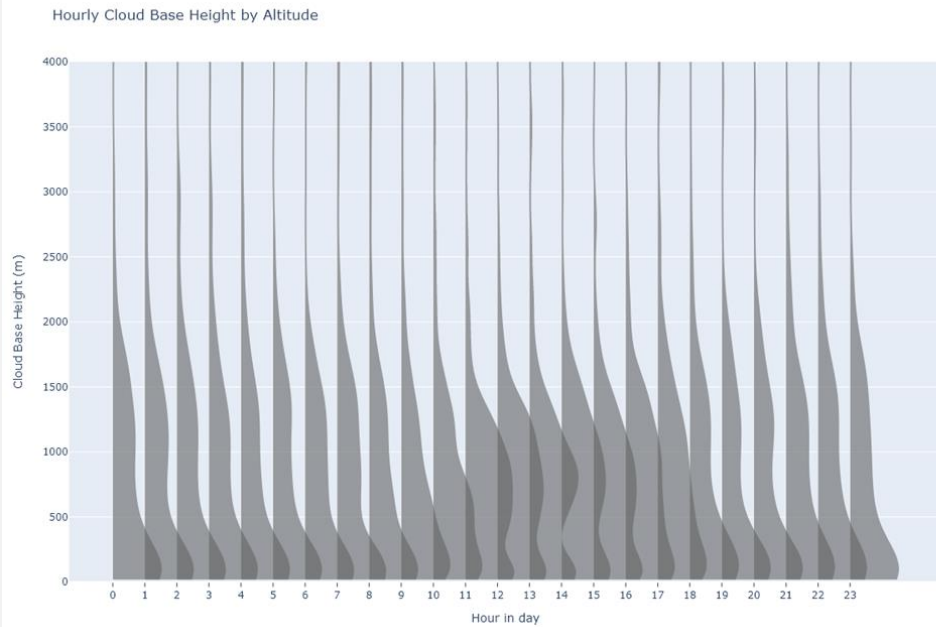
7 References

- Bodas-Salcedo, A., Andrews, T., Karmalkar, A. V., & Ringer, M. A. (2016). Cloud liquid water path and radiative feedbacks over the Southern Ocean. *Geophysical Research Letters*, *43*(20), 10,938–10,946. <https://doi.org/10.1002/2016GL070770>
- Guyot, A., Protat, A., Alexander, S. P., Klekociuk, A. R., Kuma, P., & McDonald, A. (2022). Detection of supercooled liquid water containing clouds with ceilometers: Development and evaluation of deterministic and data-driven retrievals. *Atmospheric Measurement Techniques*, *15*(12), 3663–3681. <https://doi.org/10.5194/amt-15-3663-2022>
- Kay, J. E., L'Ecuyer, T., Chepfer, H., Loeb, N., Morrison, A., & Cesana, G. (2016). Recent Advances in Arctic Cloud and Climate Research. *Current Climate Change Reports*, *2*(4), 159–169. <https://doi.org/10.1007/s40641-016-0051-9>
- Kuma, P. (2020). *Comparing remotely sensed observations of clouds and aerosols in the Southern Ocean with climate model simulations*. Zenodo. <https://doi.org/10.5281/zenodo.3865850>
- Kuma, P. (2023). *Cl2nc* [Python]. <https://github.com/peterkuma/cl2nc> (Original work published 2017)
- Kuma, P., Bender, F. A. -M., & Jönsson, A. R. (2023). Climate Model Code Genealogy and Its Relation to Climate Feedbacks and Sensitivity. *Journal of Advances in Modeling Earth Systems*, *15*(7), e2022MS003588. <https://doi.org/10.1029/2022MS003588>
- Kuma, P., McDonald, A. J., Morgenstern, O., Querel, R., Silber, I., & Flynn, C. J. (2021). Ground-based lidar processing and simulator framework for comparing models and observations (ALCF 1.0). *Geoscientific Model Development*, *14*(1), 43–72. <https://doi.org/10.5194/gmd-14-43-2021>
- Listowski, C., Delanoë, J., Kirchgaessner, A., Lachlan-Cope, T., & King, J. (2019). Antarctic clouds, supercooled liquid water and mixed phase, investigated with DARDAR: Geographical and seasonal variations. *Atmospheric Chemistry and Physics*, *19*(10), 6771–6808. <https://doi.org/10.5194/acp-19-6771-2019>
- Schuddeboom, A. J., & McDonald, A. J. (2021). The Southern Ocean Radiative Bias, Cloud Compensating Errors, and Equilibrium Climate Sensitivity in CMIP6 Models. *Journal of Geophysical Research: Atmospheres*, *126*(22), e2021JD035310. <https://doi.org/10.1029/2021JD035310>
- Trenberth, K. E., & Fasullo, J. T. (2010). Tracking Earth's Energy. *Science*, *328*(5976), 316–317. <https://doi.org/10.1126/science.1187272>
- Whitehead, L. E., McDonald, A. J., & Guyot, A. (2023). *Supercooled liquid water cloud classification using lidar backscatter peak properties* [Preprint]. Clouds/Remote Sensing/Data Processing and Information Retrieval. <https://doi.org/10.5194/egusphere-2023-1085>
- Zelinka, M. D., Myers, T. A., McCoy, D. T., Po-Chedley, S., Caldwell, P. M., Ceppi, P., Klein, S. A., & Taylor, K. E. (2020). Causes of Higher Climate Sensitivity in CMIP6 Models. *Geophysical Research Letters*, *47*(1), e2019GL085782. <https://doi.org/10.1029/2019GL085782>

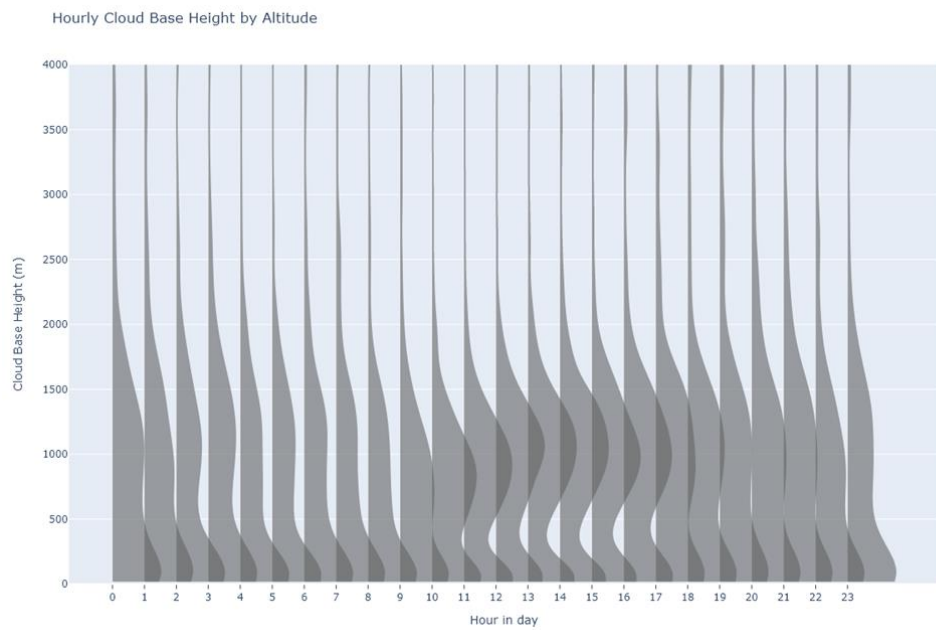
Appendix A Cloud occurrence as a function of altitude by hour

SITE

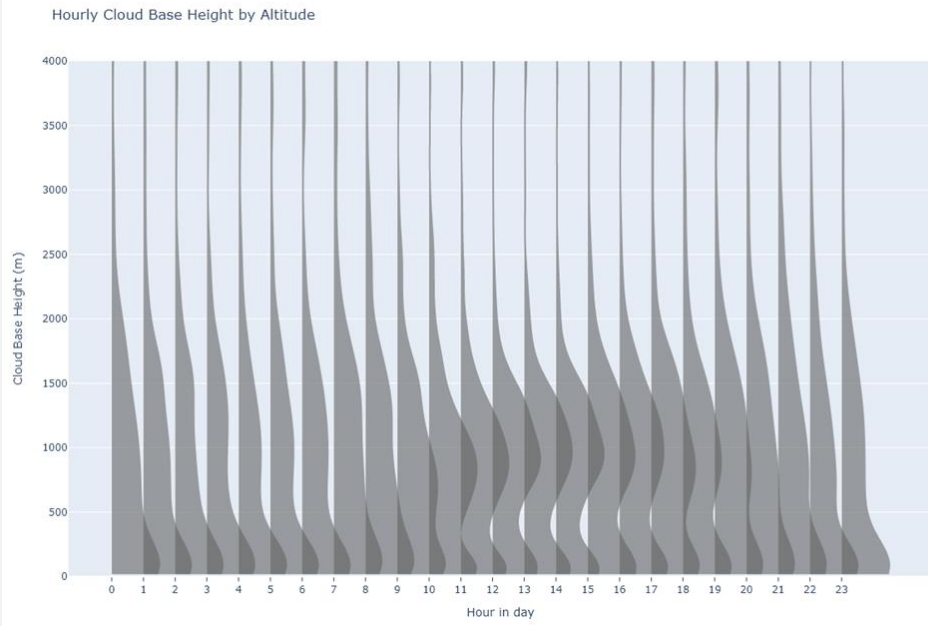
Kaitaia Aero
(NZKTA)



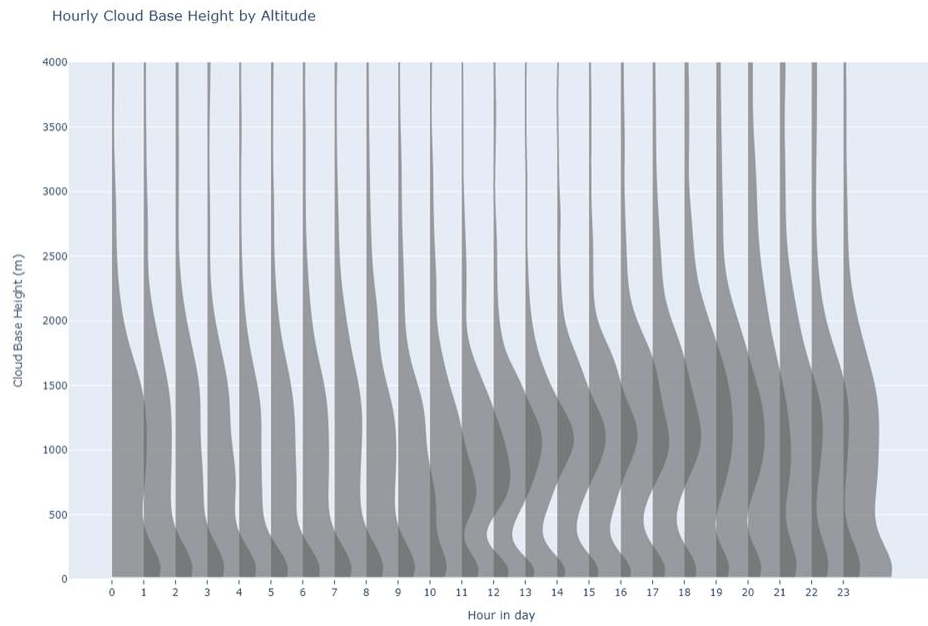
Whangarei
Aero
(NZWRA)



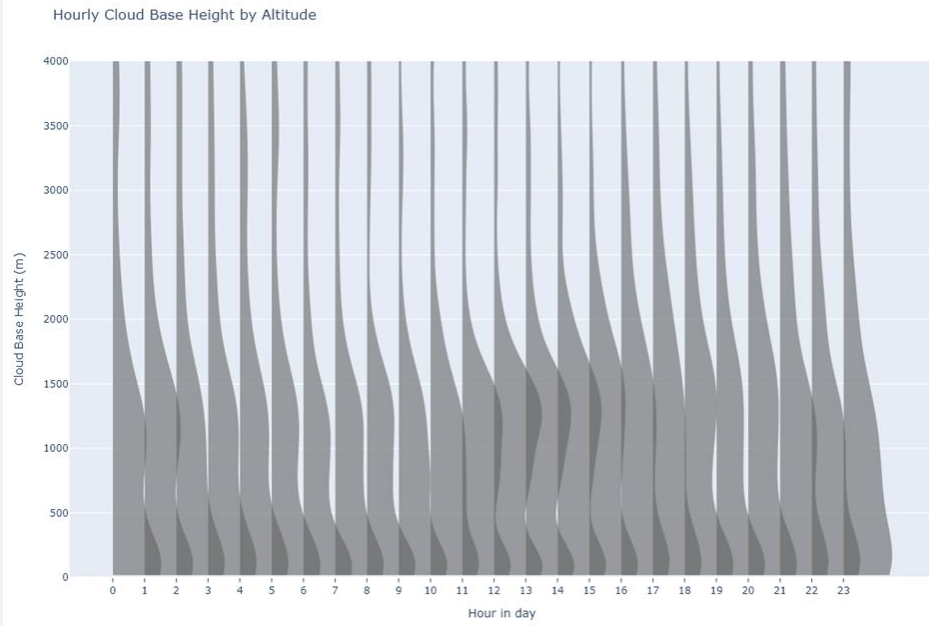
**Auckland
Aero
(NZAAA)**



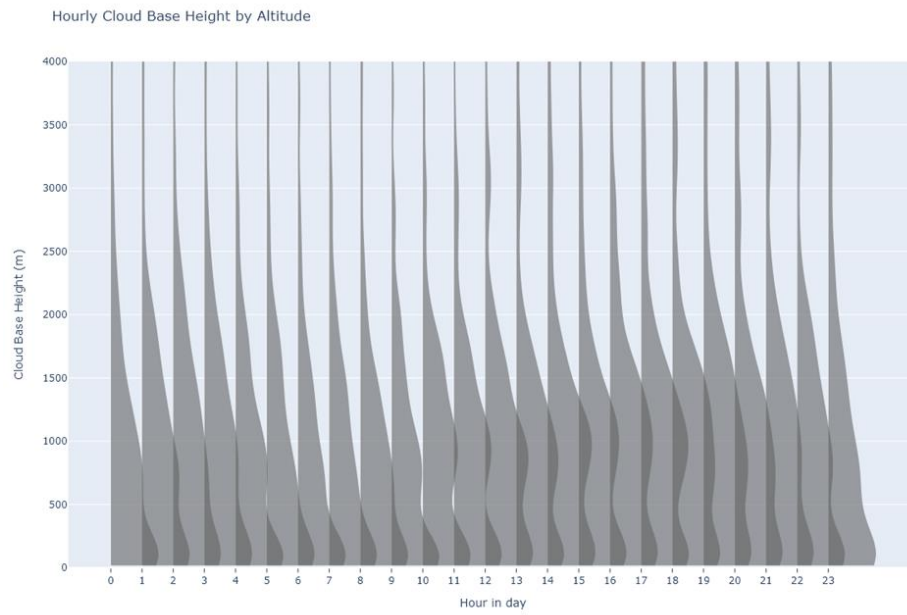
**Hamilton
Aero
(NZHNA)**



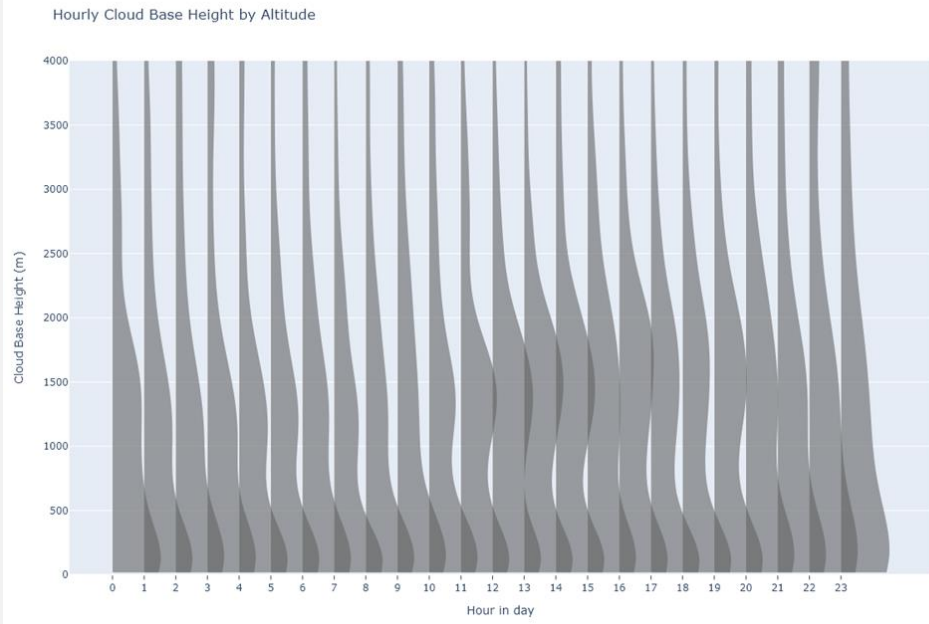
**Gisborne
Aero
(NZGSA)**



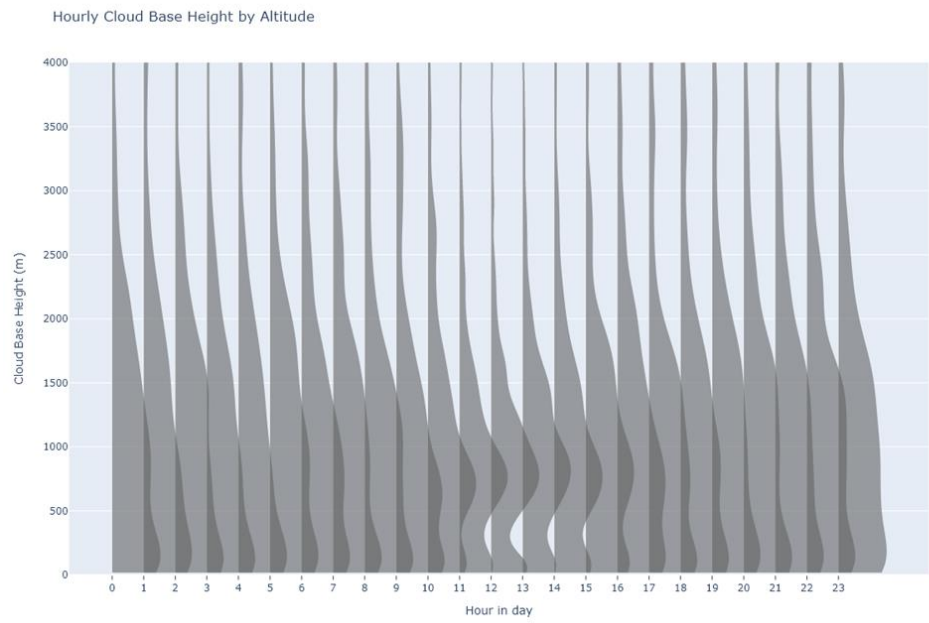
**New
Plymouth
Aero
(NZNPA)**



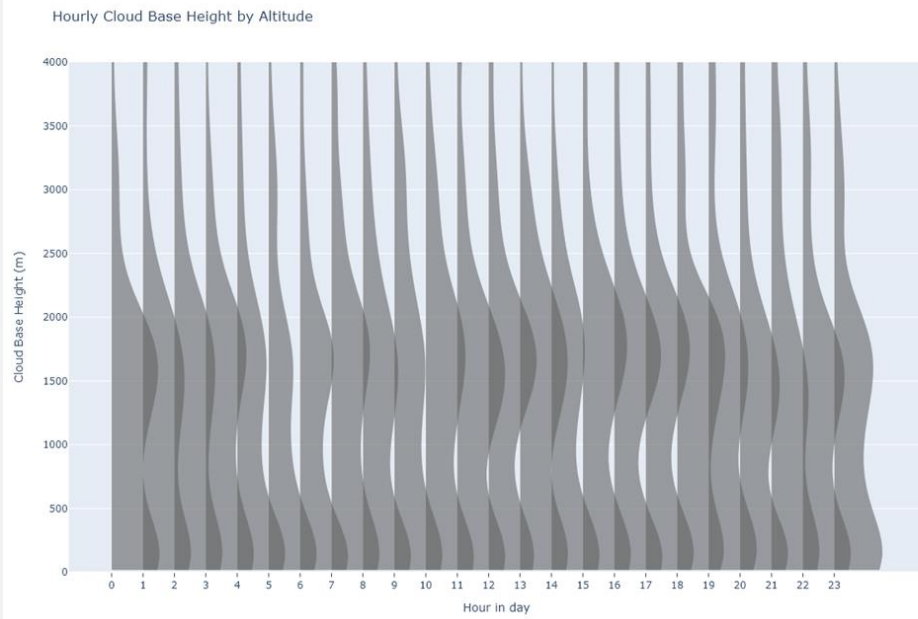
**Napier Aero
(NZNRA)**



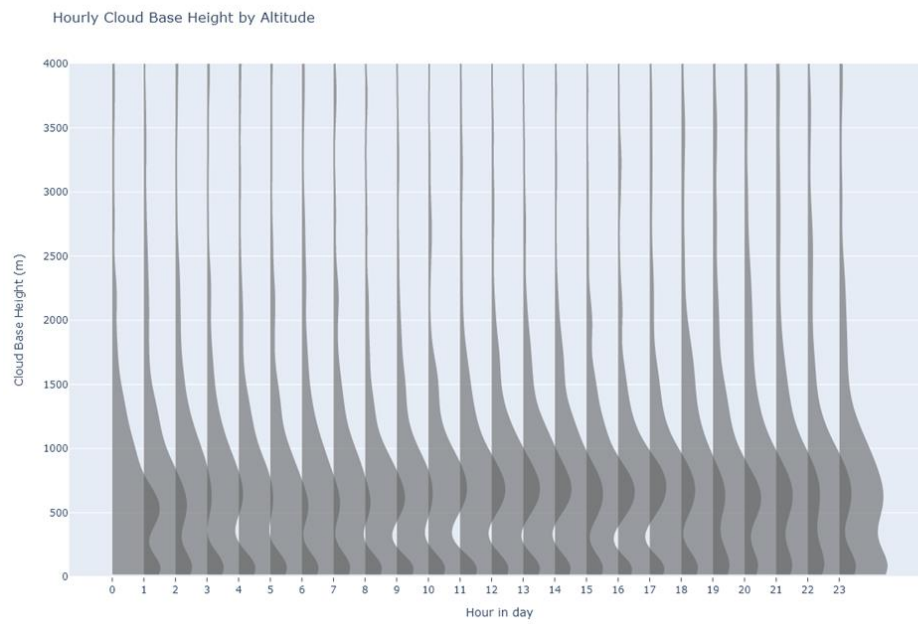
**Ohakea Aero
(NZOHA)**



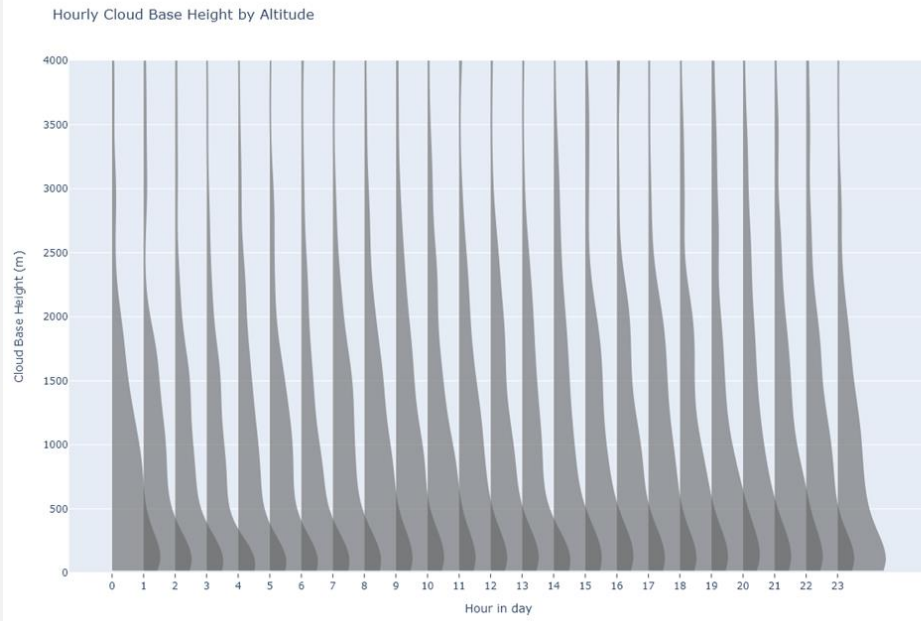
**Nelson Aero
(NZNSA)**



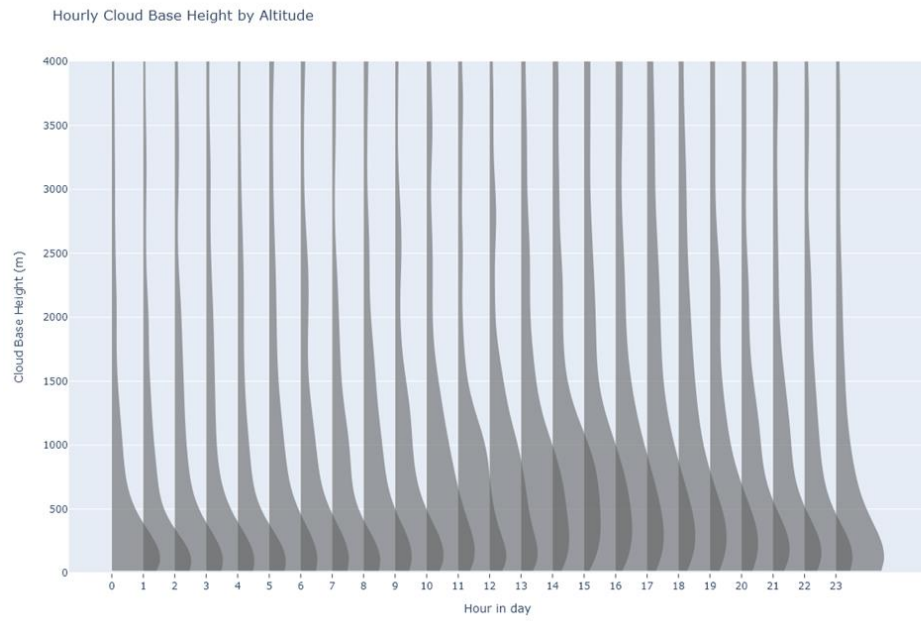
**Wellington Aero
(NZWNA)**



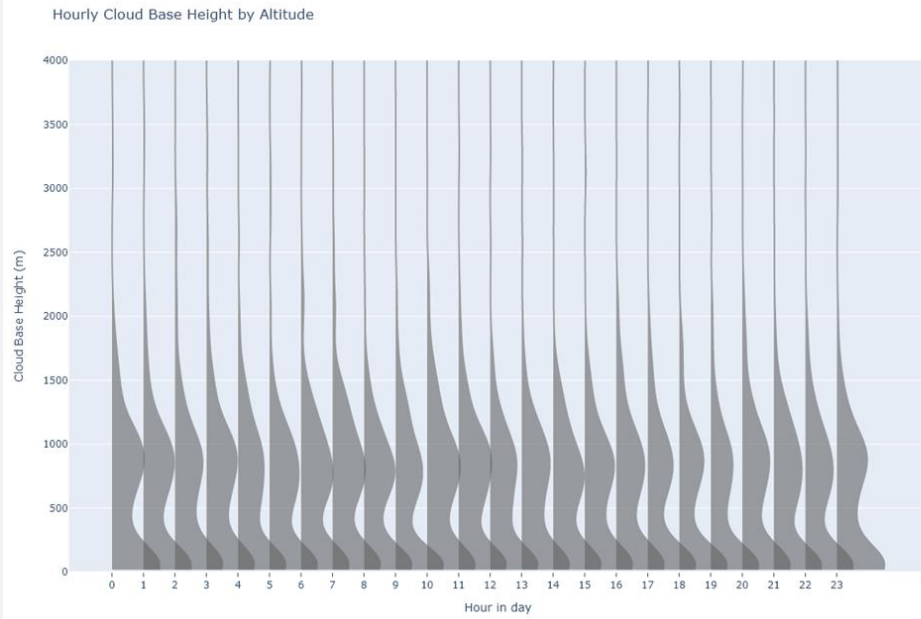
Hokitika
Aero
(NZHKA)



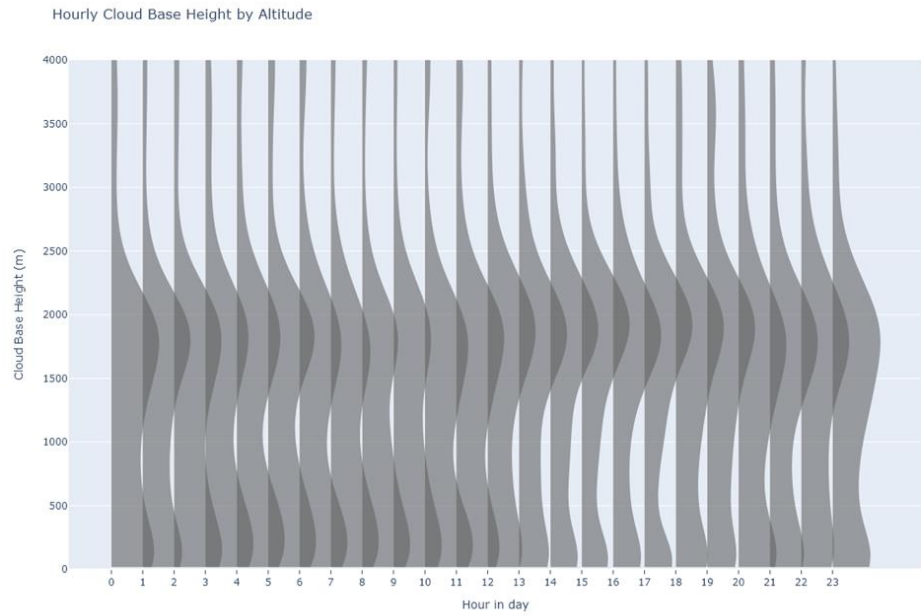
Christchurch
Aero
(NZCHA)



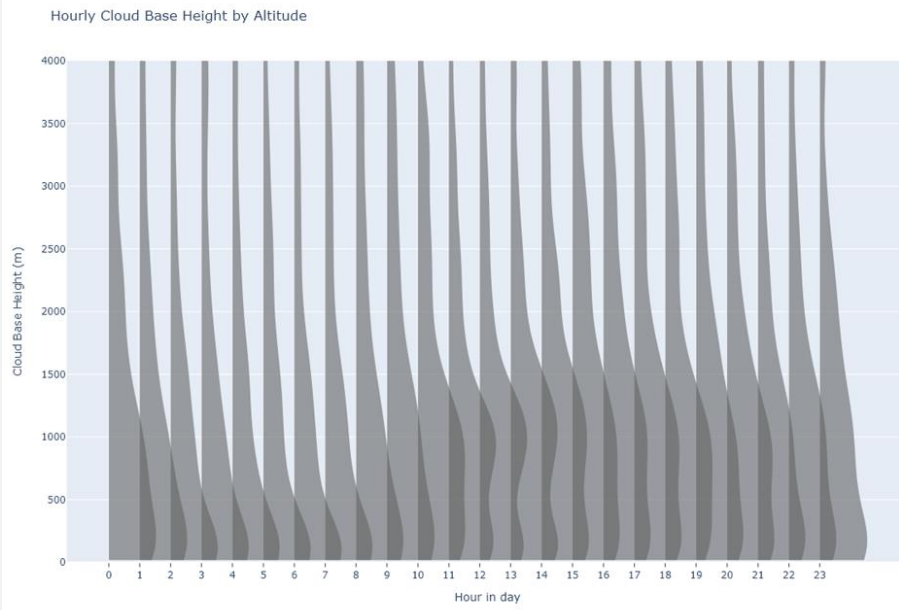
Chatham
Island Aero
(NZCIA)



Queenstown
Aero
(NZQNA)



**Dunedin
Aero
(NZDNA)**



**Invercargill
Aero
(NZNVA)**

